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Article

## Comparison of Geophysical Model Functions for SAR Wind Speed Retrieval in Japanese Coastal Waters

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**Abstract:** This work discusses the accuracies of geophysical model functions (GMFs) for retrieval of sea surface wind speed from satellite-borne Synthetic Aperture Radar (SAR) images in Japanese coastal waters characterized by short fetches and variable atmospheric stability conditions. *In situ* observations from two validation sites, Hiratsuka and Shirahama, are used for comparison of the retrieved sea surface wind speeds using CMOD (C-band model)4, CMOD\_IFR2, CMOD5 and CMOD5.N. Of all the geophysical model functions (GMFs), the latest C-band GMF, CMOD5.N, has the smallest bias and root mean square error at both sites. All of the GMFs exhibit a negative bias in the retrieved wind speed. In order to understand the reason for this bias, all SAR-retrieved wind speeds are separated into two categories: onshore wind (blowing from sea to land) and offshore wind (blowing from land to sea). Only offshore winds were found to exhibit the large negative bias, and short fetches from the coastline may be a possible reason for this. Moreover, it is clarified that in both the unstable and stable conditions, CMOD5.N has atmospheric stability effectiveness, and can keep the same accuracy with CMOD5 in the neutral condition. In short, at the moment, CMOD5.N is thought to be the most promising GMF for the SAR wind speed retrieval with the atmospheric stability correction in Japanese

coastal waters, although there is ample room for future improvement for the effect from short fetch.

**Keywords:** satellite-borne SAR; geophysical model function; sea surface wind speed retrieval

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## 1. Introduction

Estimation of sea surface wind speed has been attempted since the 1980s using a microwave scatterometer and a geophysical model function (GMF), which describes the relationship between microwave backscatter, local incidence angle, relative wind direction and wind speed at 10 m height above the mean sea level (MSL). Nowadays, wind field maps using the GMF are practically utilized [1].

Multiple GMFs with the common name CMOD (C-band model) have been developed. An early C-band GMF, CMOD2, was formulated as a prelaunch model for the ERS (European Remote Sensing Satellite)-1 based on airborne scatterometer data. CMOD4 [2] was developed based on both the ERS-1 scatterometer and Synthetic Aperture Radar (SAR) data. Although CMOD4 has been frequently validated in previous studies [3–5], Stoffelen [6] suggested that CMOD4-retrieved wind speeds are negatively biased by 4%. In addition, CMOD4 cannot retrieve high wind speeds (24 m/s and above). Meanwhile, CMOD\_IFR2 [7] was developed independently from CMOD4. CMOD\_IFR2 was also calibrated for the ERS-1 scatterometer based on buoy and analysis data. CMOD5 was developed to overcome the weak points of CMOD4, and has been used recently [8], but it still has negative biases [9]. In order to eliminate this negative bias, the latest GMF, CMOD5.N [9,10], was developed by refitting 28 coefficients of CMOD5 to reduce the negative bias, and CMOD5.N retrieves 0.5 m/s higher wind speed than CMOD5 on average for the correction.

A distinguishing feature of CMOD5.N is that it can retrieve the equivalent neutral wind speed (ENW) [11], whereas CMOD4 and CMOD5 retrieve a non-neutral wind speed, which is referred to in this study as the stability-dependent wind speed (SDW) meaning a real wind speed. A scatterometer and SAR do not directly measure the 10 m-height wind speed, but they observe backscatter from the sea surface. The intensity of the backscatter depends on the sea surface roughness or frictional velocity, which can be related to the 10 m-height wind speed through the Monin-Obukhov similarity theory [12]. That is, in the surface layer, wind speed increases with height, and the deviation from the logarithmic profile is determined by atmospheric stability. Thus, ideally, the 10 m-height wind speed should be retrieved taking atmospheric stability into account, and this is possible using CMOD5.N. The CMOD5.N-retrieved ENW can be converted to a SDW taking the atmospheric stability into account by using external input data of sea surface temperature, air temperature and relative humidity. As an offset for atmospheric stability, a value of 0.2 m/s is added in CMOD5.N compared to CMOD5 [9] on average. Along with the 0.5 m/s enhancement applied to eliminate the negative bias mentioned above, the difference between the wind speeds retrieved by CMOD5.N and CMOD5 is 0.7 m/s on average. Meanwhile, CMOD\_IFR2 also was adjusted to near neutral wind speeds [7].

The purpose of this study is to identify the most promising GMF for the SAR wind speed retrieval in Japanese coastal waters, which are greatly affected by complex coastal topography and variable atmospheric stability due to prevailing monsoon winds and warm and cold ocean currents. Four

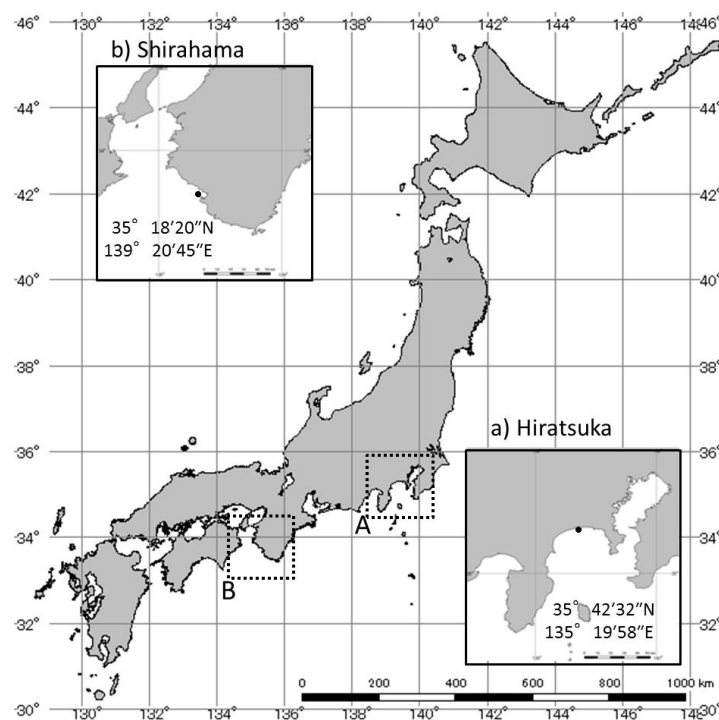
C-band GMFs; CMOD4, CMOD\_IFR2, CMOD5 and CMOD5.N, are compared with each other using 106 ENVISAT (ENVironmental SATellite) Advanced SAR images at two validation sites, Hiratsuka and Shirahama, in Japanese coastal waters. The methods are described in Section 2. General results are shown in Section 3.1, and discussions on effect of fetch, and the effectiveness of the correction for atmospheric stability with CMOD5.N are described in Sections 3.2 and 3.3, respectively. Finally, conclusions are given in Section 4.

## 2. Methods

### 2.1. In situ Measurements and Synthetic Aperture Radar (SAR) Images

The two target coastal waters of this study are Hiratsuka and Shirahama in Japan, where there are offshore platforms conducting meteorological observations. The geographical locations of the Hiratsuka offshore platform ( $35^{\circ}18'20''\text{N}$ ,  $139^{\circ}20'45''\text{E}$ ) operated by the Institute of Industrial Science (IIS) of the University of Tokyo, and the Shirahama offshore platform ( $33^{\circ}42'32''\text{N}$ ,  $135^{\circ}19'58''\text{E}$ ) operated by the Disaster Prevention Research Institute (DPRI) of Kyoto University are shown in Figure 1. These platforms have distances of 1 km and 2 km from coast lines, respectively. On the two platforms, 10-min-average wind speed and direction are measured at a height of 23 m above the MSL by a propeller anemometer. These measured wind speeds are used to validate the sea surface wind speeds retrieved from SAR images.

**Figure 1.** Geographical locations of the Hiratsuka (a) and Shirahama (b) offshore platforms (indicated by black circles in the inserts).



This study uses images from the C-band Advanced SAR onboard the ENVISAT satellite, launched by the European Space Agency in 2002. In total, 106 Advanced SAR images with HH polarization for a horizontal transmit and a horizontal receive, and VV polarization for a vertical transmit and a vertical

receive (Hiratsuka: 33 images, and Shirahama: 73 images) recorded with 12.5 m pixel spacing for the Precision Image Product (IMP) and 75.0 m pixel spacing for Wide Swath Mode (WSM) are used to retrieve sea surface wind speeds. All 106 images are listed in Appendixes 1 (Hiratsuka) and 2 (Shirahama). The SAR images are smoothed to 400 m spatial resolution using the Cressman method [13] to remove speckle noise. Though it is expected that these platforms themselves make errors in SAR images, these errors also can be eliminated by the smoothing as well as the speckle noise as previous study reported [14].

## 2.2. C-Band Geophysical Model Function (GMF)

Winds retrieved from four C-band GMFs, CMOD4, CMOD5, CMOD\_IFR2, and CMOD5.N are compared with *in situ* measurements at Hiratsuka and Shirahama. The lower limit of wind speed retrieved from a GMF is known to be 1 to 3 m/s [15–17]. Thus, this paper excludes *in situ* wind speeds of less than 2 m/s at 23 m height from the validation. For CMOD4, the primary equation is written as

$$\sigma_{vv}^0 = b_0(1.0 + b_1 \cos \phi + b_3 \tanh(b_2) \cos(2\phi))^{1.6}, \quad (1)$$

for CMOD\_IFR2, it is

$$\sigma_{vv}^0 = 10.0^{\alpha + \beta \sqrt{v}} (1.0 + b_1 \cos \phi + \tanh b_2 \cos 2\phi) \quad (2)$$

and for CMOD5 and CMOD5.N, it is

$$\sigma_{vv}^0 = b_0(1.0 + b_1 \cos \phi + b_2 \cos(2\phi))^{1.6}, \quad (3)$$

where  $\sigma_{vv}^0$  is the VV-polarized Normalized Radar Cross Section (NRCS) and  $\phi$  is the relative direction between the radar look direction and the wind direction. The other parameters,  $b_i$  ( $i = 0, 1, 2, 3$ ) depend on the radar incidence angle and wind speed. The relative wind direction should be acquired from other external sources, whereas the NRCS and the incidence angle can be obtained from an SAR image. *In situ* wind directions are used as input to the GMFs in this study, although some methods to acquire wind directions from an SAR image itself or numerical simulations have been proposed (e.g., [18–20]).

Since all the C-band GMFs have been developed for VV-polarized NRCS ( $\sigma_{vv}^0$ ), HH-polarized NRCS ( $\sigma_{hh}^0$ ) must be corrected by an empirical equation before the GMF processing. In order to convert from  $\sigma_{hh}^0$  to  $\sigma_{vv}^0$ , the equation [21]:

$$\frac{\sigma_{vv}^0}{\sigma_{hh}^0} = C_0(\theta) + C_1(\theta) \cos \phi + C_2(\theta) \cos 2\phi \quad (4)$$

is applied.  $C_i$  ( $i = 0, 1, 2$ ) are the ratios between HH and VV-polarized NRCSs for three wind directions, upwind, downwind and crosswind, respectively.

## 2.3. Height Correction of In situ Wind Speed

A GMF can retrieve the sea surface wind speed at a 10 m height above the MSL, whereas the *in situ* wind speed is measured at 23 m above MSL at both offshore platforms. Thus, height correction of the *in situ* wind speed is necessary to enable direct comparison with the retrieved wind speeds from the SAR images. For this height correction, the LKB code [22] is used to estimate the 10 m-height wind speed from 23 m-height wind speed. The LKB code requires three kinds of input data, air

temperature, relative humidity, and sea surface temperature (SST) for the estimation. This study principally uses values from *in situ* measurements. But, since air temperature is not measured at Hiratsuka, those simulated with the meteorological mesoscale model WRF (the Weather Research and Forecasting model [23]) are used instead of *in situ* measurements. For this simulation, the Advanced Research WRF model version 3.0 is used, and the model configuration is shown in Appendix. In the LKB code, the following wind profile based on the Monin-Obukhov similarity theory is used, taking the effect of atmospheric stability, expressed as  $\Psi u(z/L)$ , into account.

$$u = \frac{u_*}{\kappa} \left[ \ln \left( \frac{z}{z_0} \right) - \Psi_u \left( \frac{z}{L} \right) \right] \quad (5)$$

Here,  $u_*$  is frictional velocity,  $z_0$  is roughness length,  $z$  is height,  $L$  is Monin-Obukhov length, and  $\kappa$  is the von Karman constant ( $=0.4$ ). The relation between  $z_0$  and  $u_*$  is given as

$$z_0 = 0.11 \frac{\nu}{u_*} + z_{ch} \frac{u_*^2}{g} \quad (6)$$

where  $z_{ch}$  is Charnock's parameter with a value of 0.011 [24],  $\nu$  is the kinematic viscosity, and  $g$  is the acceleration due to gravity. The parameters of  $z_0$  and  $u_*$  can be determined iteratively.

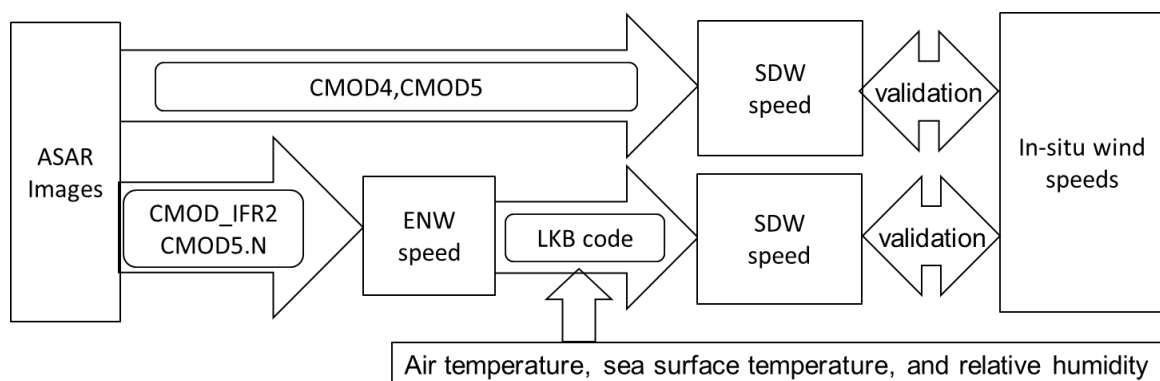
#### 2.4. Conversion from ENW to SDW

ENW represents the wind at 10 m height for given surface stress in case the marine boundary layer is neutrally stratified, while SDW means real wind speed. In order to compare CMOD\_IFR2- and CMOD5.N-retrieved ENW with *in situ* wind speeds in the validation, the ENW is converted to SDW, which is comparable to the *in situ* wind speed. Hereinafter, the SDW obtained from CMOD\_IFR2 and CMOD5.N with the LKB code is referred to as CMOD\_IFR2\_SDW and CMOD5.N\_SDW, while the ENW originally obtained from CMOD\_IFR2 and CMOD5.N is referred to as CMOD\_IFR2\_ENW and CMOD5.N\_ENW. The logarithmic wind profile,

$$u_{ENW} = \frac{u_*}{\kappa} \ln \left( \frac{z}{z_0} \right) \quad (7)$$

is first used for the calculation of frictional velocity  $u_*$  from the ENW. Then the SDW is calculated using Equation (5). The flowchart of the wind retrieval from an Advanced SAR image and its validation with an *in situ* wind speed is depicted in Figure 2.

**Figure 2.** Flowchart of wind retrieval from an Advanced SAR image and validation with *in situ* wind speed.

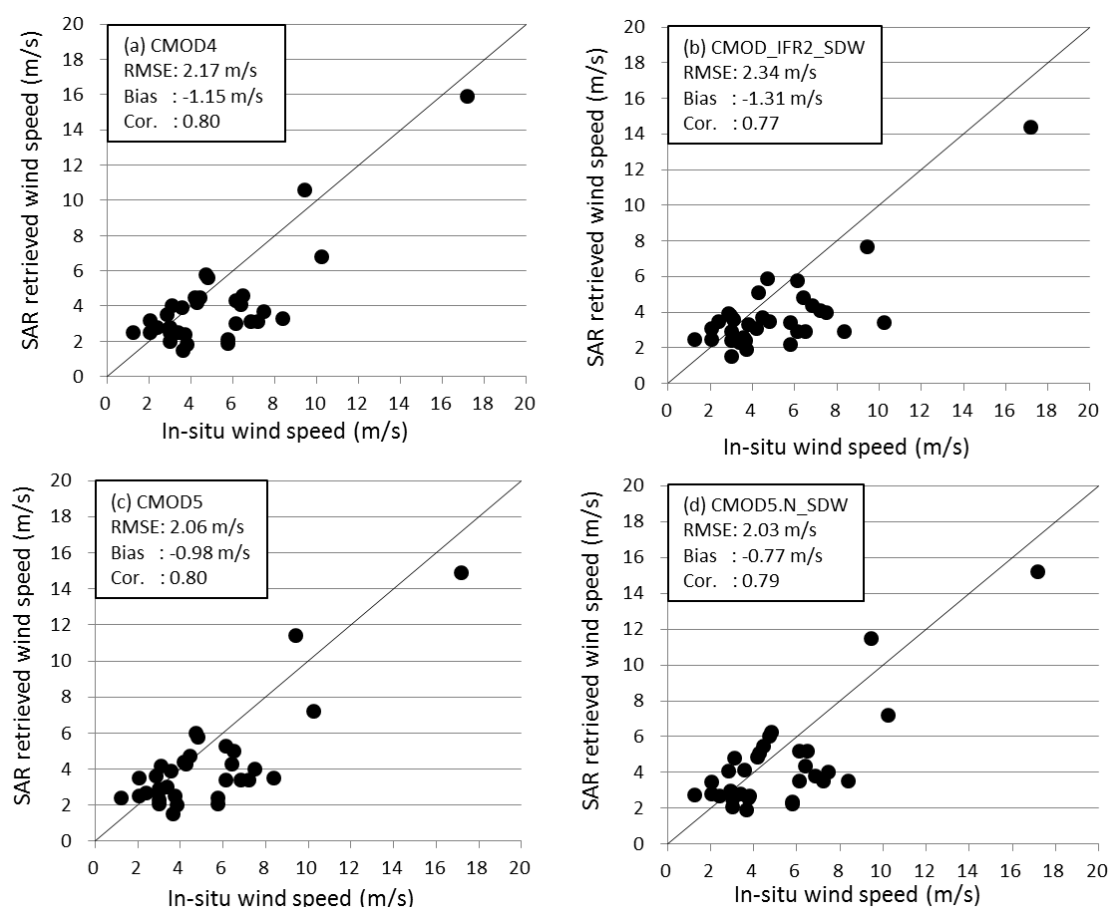


### 3. Results and Discussion

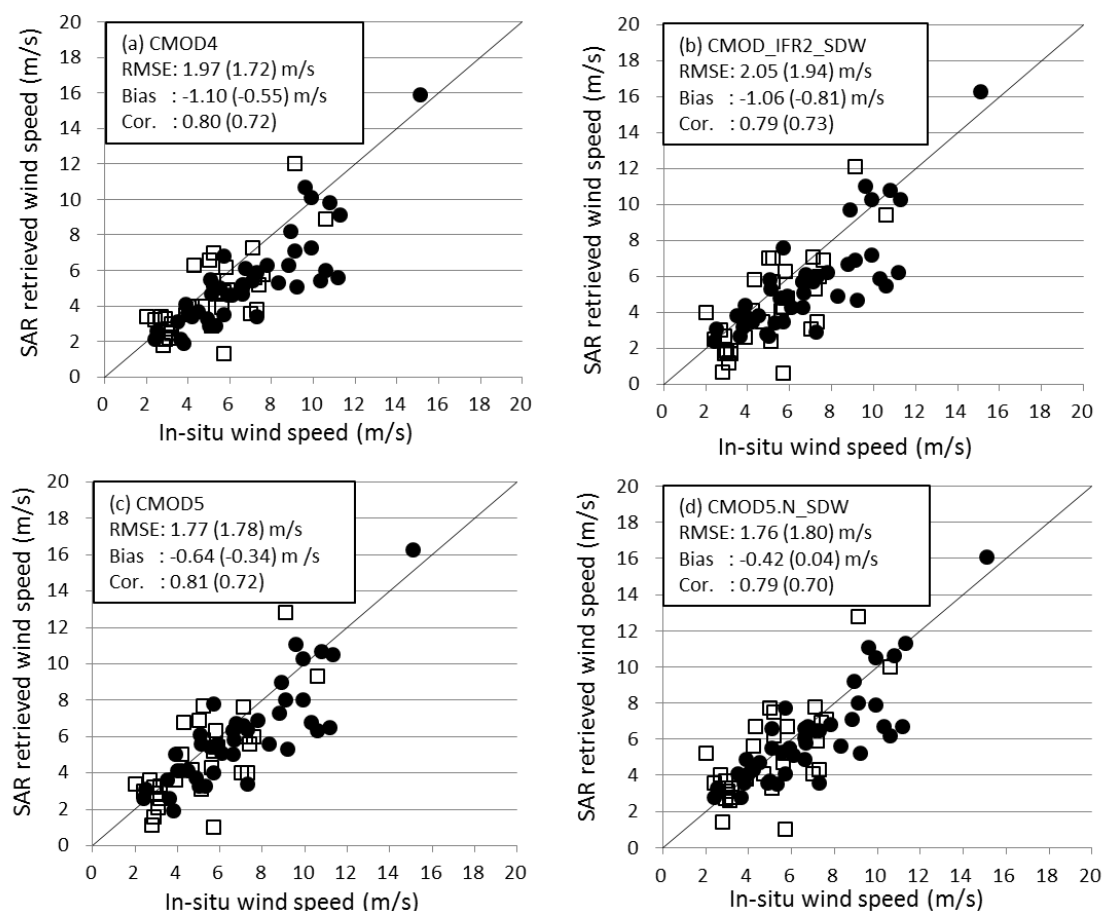
#### 3.1. Accuracies of Wind Speed Retrieval Using Four GMFs

Accuracies of the wind speed retrieved by four GMFs: CMOD4, CMOD\_IFR2, CMOD5, and CMOD5.N, are compared at the two validation sites, Hiratsuka and Shirahama. Figure 3 shows relations between *in situ* wind speeds and retrieved wind speeds from 33 SAR images using the four GMFs at Hiratsuka. The root mean square errors (RMSEs) of these SAR-retrieved wind speeds are 2.17 m/s (CMOD4), 2.34 m/s (CMOD\_IFR2\_SDW), 2.06 m/s (CMOD5) and 2.03 m/s (CMOD5.N\_SDW), and the biases are −1.15 m/s (CMOD4), −1.31 m/s (CMOD\_IFR2\_SDW), −0.98 m/s (CMOD5) and −0.77 m/s (CMOD5.N\_SDW), respectively. Thus, CMOD5.N\_SDW has the smallest RMSE and bias at Hiratsuka. Figure 4 is the same as Figure 3, but for Shirahama, using 73 SAR images. The RMSEs of these SAR-retrieved wind speeds are 1.97 m/s (CMOD4), 2.05 m/s (CMOD\_IFR2\_SDW), 1.77 m/s (CMOD5) and 1.76 m/s (CMOD5.N\_SDW), respectively and the biases are −1.10 m/s (CMOD4), −1.06 m/s (CMOD\_IFR2\_SDW), −0.64 m/s (CMOD5) and −0.42 m/s (CMOD5.N\_SDW), respectively. All the statistic values are concluded in Table 1. These results show that CMOD5.N\_SDW has the smallest RMSE and bias at Shirahama as well as Hiratsuka. Meanwhile, the retrieved wind speeds from WSM are also shown as squares in Figure 4, and the statistic values are shown in parentheses. No large difference is found in wind speed between WSM and IMP for all the GMFs.

**Figure 3.** Relations between 33 SAR-retrieved wind speeds and *in situ* wind speeds at Hiratsuka using four GMFs: (a) CMOD4, (b) CMOD\_IFR2, (c) CMOD5, and (d) CMOD5.N.



**Figure 4.** Same as Figure 3, but for 73 cases observed by IMP (daub circle) and WSM (square) at Shirahama. The statistics are for 73 cases, and those for WSM (31 cases) are in parentheses.



**Table 1.** RMSEs, biases and correlation coefficients of wind speeds retrieved by four GMFs at both Hiratsuka and Shirahama.

		CMOD4	CMOD_IFR2_SDW	CMOD5	CMOD5.N_SDW
Hiratsuka	RMSE (m/s)	2.17	2.34	2.06	2.03
	Bias (m/s)	-1.15	-1.31	-0.98	-0.77
	Correlation coef.	0.80	0.77	0.80	0.79
Shirahama	RMSE (m/s)	1.97	2.05	1.77	1.76
	Bias (m/s)	-1.10	-1.06	-0.64	-0.42
	Correlation coef.	0.80	0.79	0.81	0.79

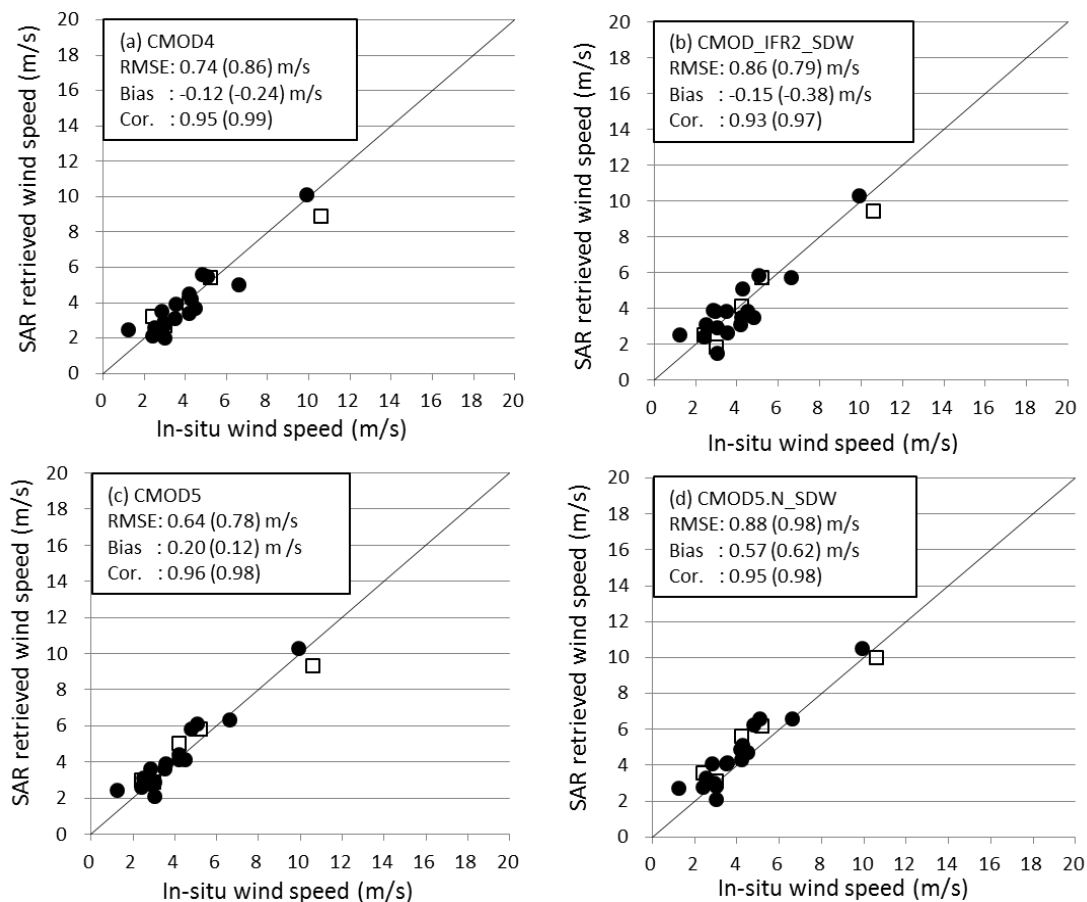
All of the GMFs exhibit a negative bias in the retrieved wind speed at both validation sites. The absolute value of the negative bias in CMOD5.N\_SDW is the smallest of all the GMFs, followed by CMOD5, CMOD4, and CMOD\_IFR2\_SDW in ascending order at both validation sites. These negative biases from CMOD4 and CMOD5 were reported in the previous study [10] as well. However, CMOD5.N, which has been developed to eliminate the negative bias, is found to still have a negative bias in this study. Moreover, the negative bias is much larger than that reported in previous studies such as [25,26], which showed that even the negative bias in CMOD5 is around -0.5 m/s. These facts indicate that the SAR wind speeds retrieved at Hiratsuka and Shirahama are affected by some factors that cause the larger negative biases. A possible factor for the negative bias is the effect of a short fetch as described in the previous study [10]. This short fetch effect is discussed in detail in Section 3.2.



### 3.2. Effect of Short Fetch on Wind Speed Retrieval

In addition to wind speed, the fetch and duration of wind blowing over the sea surface can also change the sea surface roughness, especially in coastal waters. Thus, in this study, a short fetch is suspected to affect the sea surface roughness, resulting in the negative bias in the retrieved wind speed. To extract the short fetch effect on the SAR-retrieved wind speed, all the data are separated into two categories: offshore wind (blowing from land to sea) and onshore wind (blowing from sea to land). The onshore wind is defined as wind with a wind direction between 110 to 210 degrees at Hiratsuka, and 205 to 310 degrees at Shirahama, based on a threshold fetch of 30 km. This is because the previous study [27] shows that the wind speed growth with fetch is influenced by upwind terrain within 30 km from the coastline. Figure 5 shows scatter plots of 22 onshore winds at both validation sites for the four GMFs. The RMSEs of the SAR-retrieved wind speeds are all below 1.0 m/s: 0.74 m/s (CMOD4), 0.86 m/s (CMOD\_IFR2\_SDW), 0.64 m/s (CMOD5) and 0.88 m/s (CMOD5.N\_SDW), and the biases are  $-0.12$  m/s (CMOD4),  $-0.15$  m/s (CMOD\_IFR2\_SDW),  $0.20$  m/s (CMOD5) and  $0.57$  m/s (CMOD5.N\_SDW). All the negative biases shown in Figure 3 and 4 are obviously reduced in the case of only onshore wind.

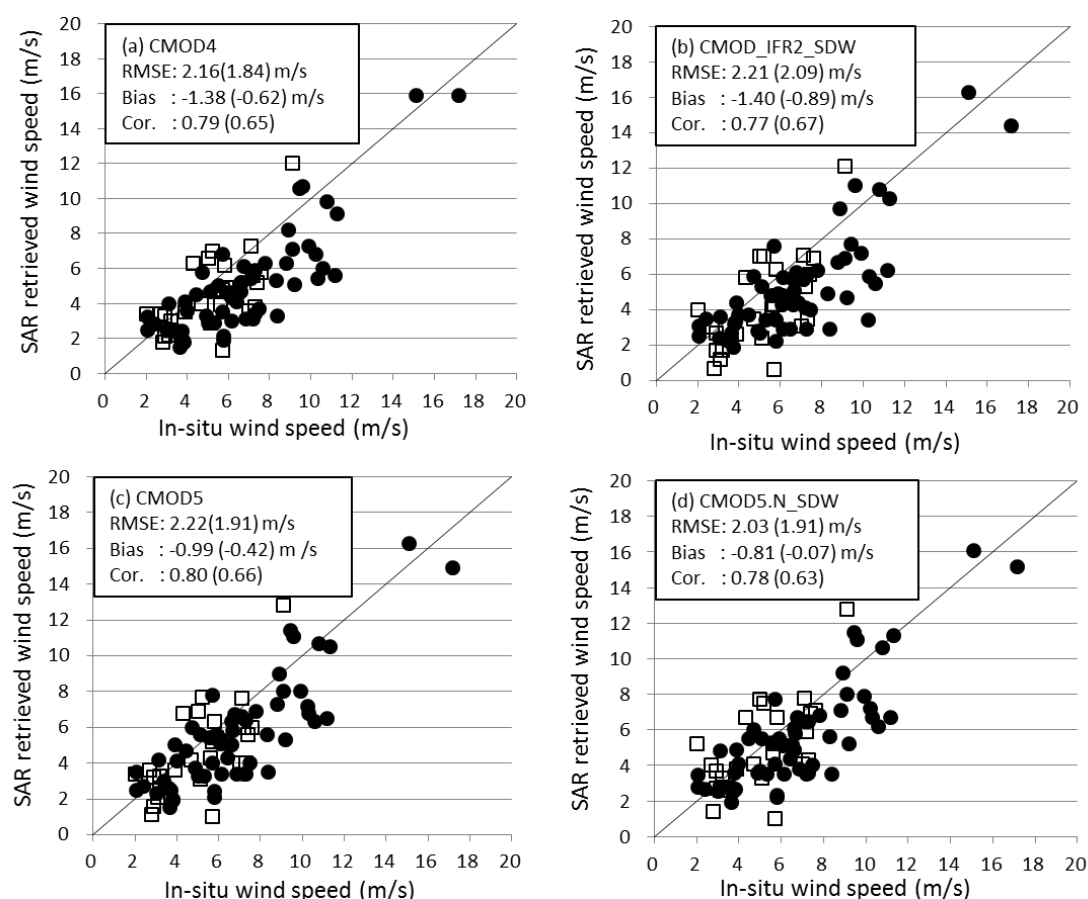
**Figure 5.** Same as Figure 3, but for 22 onshore winds by IMP (daub circle) and WSM (square) at both Hiratsuka and Shirahama. The statistics are for 22 cases, and those for WSM (5 cases) are in parentheses.



Meanwhile, Figure 6 shows the same scatter plots as Figure 5 but for 84 offshore winds. The RMSEs are 2.16 m/s (CMOD4), 2.21 m/s (CMOD\_IFR2\_SDW), 2.22 m/s (CMOD5) and 2.03 m/s

(CMOD5.N\_SDW), and the biases are  $-1.38$  m/s (CMOD4),  $-1.40$  m/s (CMOD\_IFR2\_SDW),  $-0.99$  m/s (CMOD5) and  $-0.81$  m/s (CMOD5.N\_SDW). These statistic values are concluded in Table 2. All the RMSEs are greater than 2 m/s, and are obviously larger than those in Figures 3 and 4. In addition, the negative bias is larger compared to those in Figure 3 and 4. These results demonstrate that the negative biases, shown in Figures 3 and 4, are mainly due to the short fetch effect. In addition, it is found that only the wind speeds retrieved from WSM images have smaller negative biases, compared to those from both mode images. This tendency cannot be found in the onshore wind case (Figure 5). These facts suggest that the wind speeds retrieved from the IMP image can be more easily affected by the short fetch.

**Figure 6.** Same as Figure 3, but for 84 offshore winds by IMP (daub circle) and WSM (square) at both Hiratsuka and Shirahama. The statistics are for 84 cases, and those for WSM (26 cases) are in parentheses.



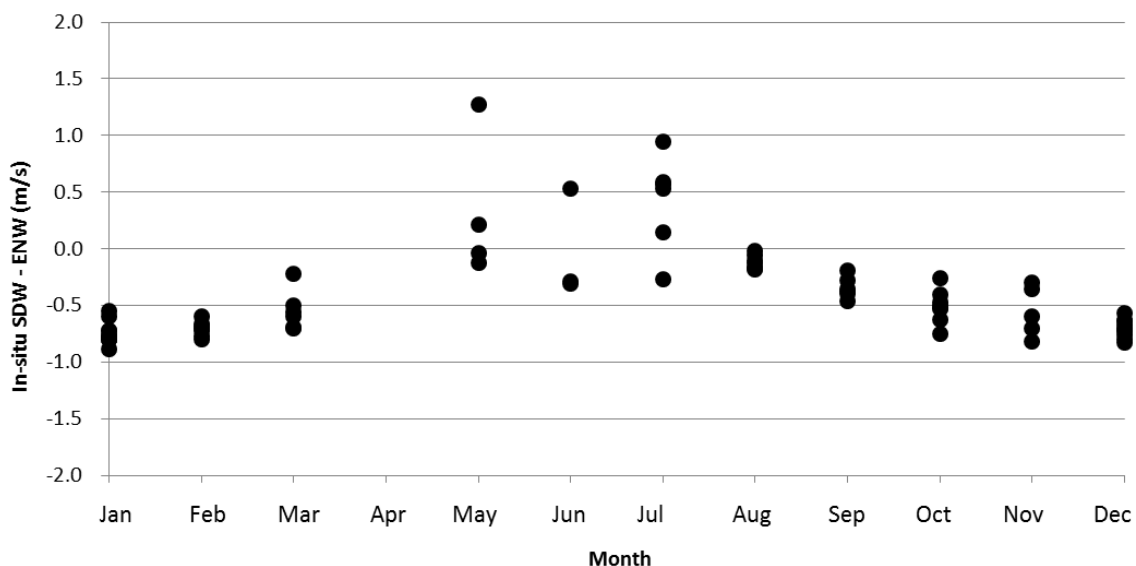
**Table 2.** RMSEs, biases and correlation coefficients of wind speeds retrieved by four GMFs at onshore and offshore winds.

		CMOD4	CMOD_IFR2_SDW	CMOD5	CMOD5.N_SDW
Onshore	RMSE (m/s)	0.74	0.86	0.64	0.88
	Bias (m/s)	-0.12	-0.15	0.20	0.57
	Correlation coef.	0.95	0.93	0.96	0.95
Offshore	RMSE (m/s)	2.16	2.21	2.22	2.03
	Bias (m/s)	-1.38	-1.40	-0.99	-0.81
	Correlation coef.	0.79	0.77	0.80	0.78

### 3.3. Effectiveness of Atmospheric Stability Correction with CMOD5.N

In this subsection, effectiveness of the atmospheric stability correction is examined. Figure 7 shows monthly differences between SDW and ENW for *in situ* measurements at Shirahama. Positive values indicate that SDW is larger than ENW. The differences between SDW and ENW range mostly between  $-0.9$  and  $-0.2$  m/s in the winter season (from October to March) when unstable conditions prevail, while the differences range from  $-0.5$  to  $1.2$  m/s in the summer season (from April to September) when neutral and stable conditions tend to occur. Roughly speaking, the differences range from  $-1$  to  $+1$  m/s throughout the year. In Figure 8, monthly variation of the stability parameter  $z/L$  is shown. The parameters do not appear unless the CMOD5.N-retrieved wind speed is 2 m/s or higher because lower wind speeds give an extremely large value of the parameter. Figure 8 additionally shows that the stability parameter also indicates the seasonal variation of the atmospheric stability. In winter season, all values of  $z/L$  range from  $-7$  to  $0$ , while  $z/L$  ranges between  $-4$  and  $1$  in summer season. From Figures 7 and 8, it is found that the difference between SDW and ENW shows seasonal variation, which depends on the actual atmospheric stability. This suggest that the atmospheric stability can cause an error of  $-1$  to  $+1$  m/s in the SAR retrieved wind speed when using the GMFs that do not take into account the effect of atmospheric stability. Thus, the atmospheric stability correction is indispensable in the SAR wind speed retrieval in Japanese coastal water.

**Figure 7.** Monthly differences (m/s) between SDW and ENW for *in situ* measurements at Shirahama.

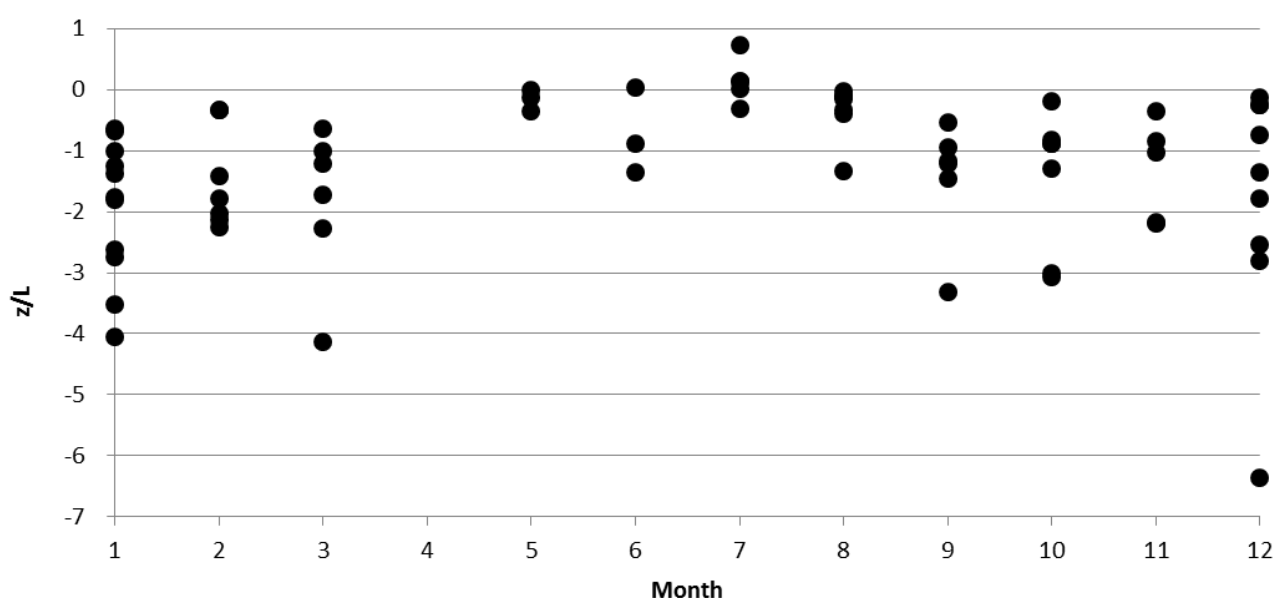


The relation between CMOD5, CMOD5.N\_ENW and CMOD5.N\_SDW is shown in Figure 9, based on the results in Section 3.1 and Figure 7. As shown in Figure 9, it is expected that the difference between CMOD5 and CMOD5.N\_SDW decreases in the unstable condition, whereas it increases in the unstable condition. The small difference in the unstable condition is caused by the correction offset of about 0.7 m/s, which was enhanced when CMOD5.N had been developed from CMOD5 as shown in Figure 9.

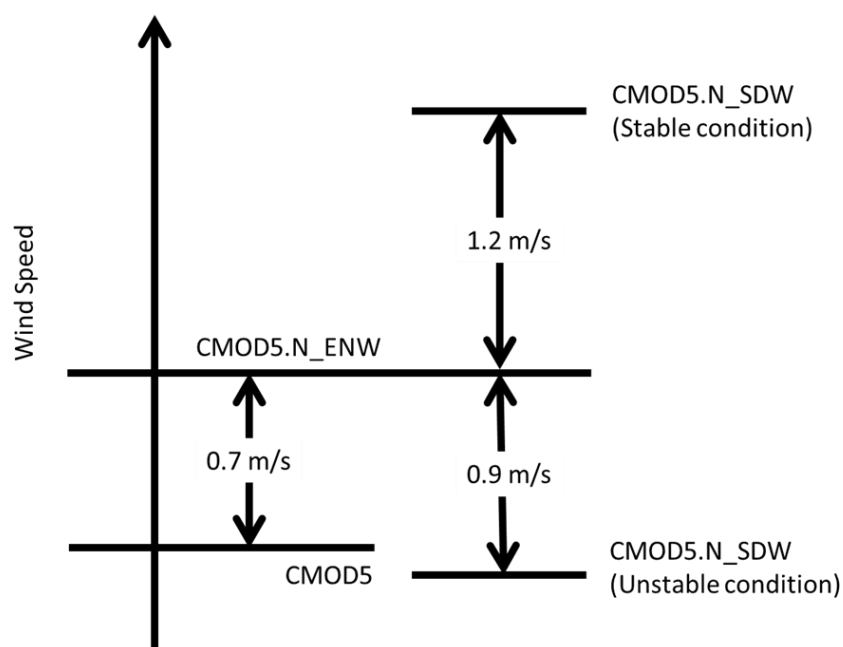
In order to examine if CMOD5.N is effective for the atmospheric stability correction, the statistical values obtained from CMOD5.N\_SDW and CMOD5 are shown in Table 3. In the neutral condition

( $-1.0 < z/L \leq 0.1$ ), the RMSE with CMOD5.N\_SDW (1.72 m/s) is slightly larger than CMOD5 (1.70 m/s). The differences of absolute biases are also few between CMOD5.N\_SDW and CMOD5. Meanwhile, in unstable ( $z/L \leq -1.0$ ) and stable condition ( $0.1 < z/L$ ), CMOD5.N\_SDW has a smaller RMSE than CMOD5. In particular, in the stable condition, the differences of RMSE and absolute biases between CMOD5.N\_SDW and CMOD5.N are 0.15 m/s and 0.68 m/s, respectively, and the difference between CMOD5.N\_SDW and CMOD5 is remarkable. These results indicate that atmospheric stability should be taken into account in the SAR wind speed retrieval in the seas with non-neutral conditions, and CMOD5.N is effective for the atmospheric stability correction.

**Figure 8.** Monthly variation of  $z/L$  at Shirahama. But only 2 m/s or higher of the CMOD5.N-retrieved wind speed.



**Figure 9.** Relationship of retrieved wind speeds among CMOD5, CMOD5.N\_ENW and CMOD5.N\_SDW under stable and unstable atmospheric conditions.



**Table 3.** Comparison of statistics between CMOD5.N\_SDW and CMOD5.

Atmospheric Condition		Unstable ( $z/L \leq -1.0$ )	Neutral ( $-1.0 < z/L \leq 0.1$ )	Stable ( $0.1 < z/L$ )
Bias (m/s)	CMOD5.N_SDW	−0.97	0.16	0.36
	CMOD5	−1.01	−0.11	−1.04
RMSE (m/s)	CMOD5.N_SDW	1.83	1.72	1.47
	CMOD5	1.85	1.70	1.62

#### 4. Conclusions

In this paper, four C-band geophysical model functions (GMFs) for sea surface wind speed retrieval: CMOD4, CMOD\_IFR2, CMOD5, and CMOD5.N, are compared using 106 SAR images at two validation sites, Hiratsuka and Shirahama, in Japanese coastal waters. The effect of the correction of atmospheric stability, which is taken into account in the latest GMF CMOD5.N, is also examined since atmospheric stability is variable in Japanese coastal waters. The main results of this study are summarized as follows.

1. Of all the GMFs, the stability-dependent wind speed (SDW) retrieved with CMOD5.N and corrected for atmospheric stability with the LKB code (CMOD5.N\_SDW) has the smallest root mean square error (RMSE) and the smallest bias at both Hiratsuka and Shirahama. The RMSEs are 2.03 m/s at Hiratsuka and 1.76 m/s at Shirahama, and the biases are −0.77 m/s at Hiratsuka and −0.42 m/s at Shirahama.
2. All of the GMFs exhibit a negative bias in the retrieved wind speed at both validation sites. By contrasting winds blowing onshore and offshore, only offshore winds were found to have a large negative bias. This indicates that the negative bias is primarily caused by short fetches from a coastline.
3. Atmospheric stability can cause an error of about −1 to +1 m/s in the SAR retrieved wind speed at Shirahama when using the GMFs that do not take the effect of atmospheric stability into account. CMOD5.N can reduce this error especially in non-neutral conditions, indicating that the use of CMOD5.N is effective for the atmospheric stability correction.

In short, at the moment, CMOD5.N is thought to be the most promising GMF for SAR wind speed retrieval in Japanese coastal waters. However, it is also clear that there is still ample room for future improvement. For instance, it is necessary to examine a larger number of wind speed samples because the samples used in this study were mostly less than 12 m/s. Future work is also necessary for the investigation of the short fetch effect causing a large negative bias in winds blowing offshore.

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## Appendix

**Appendix 1.** List of 33 ENVISAT Advanced SAR images with corresponding *in situ* measurements at Hiratsuka.

SAR				<i>In situ</i>		WRF	
Date (year/month/day)	Time (hr:min:sec)	Polarization	Observation Mode	Wind Direction ( °C)	Wind Speed (m s <sup>-1</sup> ) at 10 m	Sea Temp. ( °C)	Air Temp. ( °C)
2003/02/01	00:54:51	VV	IMP	3	6.1	17.8	8.6
2003/02/17	00:51:57	VV	IMP	350	3.7	18.1	11.6
2003/06/21	00:54:56	VV	IMP	213	3.8	23.8	25.3
2003/07/26	00:55:00	VV	IMP	5	5.8	24.3	25.5
2003/08/30	00:55:04	VV	IMP	19	3.6	26.5	28.0
2003/10/04	00:55:00	VV	IMP	38	3.4	26.0	21.0
2003/10/20	00:52:06	VV	IMP	1	8.4	24.2	18.5
2003/11/08	00:54:56	VV	IMP	117	3.0	24.4	20.9
2003/12/25	12:31:35	VV	IMP	233	17.2	20.2	15.7

## Appendix 1. Cont.

SAR				<i>In situ</i>		WRF	
Date (year/month/day)	Time (hr:min:sec)	Polarization	Observation Mode	Wind Direction ( °C)	Wind Speed (m s <sup>-1</sup> ) at 10 m	Sea Temp. ( °C)	Air Temp. ( °C)
2003/12/29	00:52:09	VV	IMP	259	9.4	20.1	13.4
2005/10/05	00:49:22	HH	IMP	40	3.0	22.6	22.0
2005/11/25	00:46:29	HH	IMP	24	4.7	22.9	14.8
2005/12/30	00:46:22	HH	IMP	18	2.4	16.9	7.2
2006/03/10	00:46:20	HH	IMP	357	10.2	17.1	10.2
2006/05/19	00:46:26	HH	IMP	17	4.4	20.5	21.5
2006/06/23	00:46:31	HH	IMP	298	2.0	22.6	22.4
2006/07/28	00:46:32	HH	IMP	168	4.3	24.4	25.4
2007/01/22	00:52:02	VV	IMP	352	7.5	18.7	9.7
2007/05/04	00:46:26	HH	IMP	121	3.6	21.6	20.2
2007/05/07	00:52:06	HH	IMP	174	3.0	21.1	19.8
2007/06/08	00:46:28	VV	IMP	146	4.2	21.7	21.2
2007/08/17	00:46:28	HH	IMP	146	2.8	29.4	30.7
2007/09/21	00:46:23	VV	IMP	183	1.2	25.9	25.7
2007/11/17	00:54:53	VV	IMP	3	5.8	22.8	13.7
2007/11/30	00:46:19	HH	IMP	353	6.4	20.2	13.3
2007/12/03	00:51:59	HH	IMP	212	7.2	19.6	15.3
2008/04/21	00:52:00	VV	IMP	8	6.9	19.3	19.7
2008/06/27	00:46:22	HH	IMP	83	3.1	21.0	22.2
2008/08/01	00:46:22	HH	IMP	165	2.9	22.2	20.9
2008/10/13	00:52:00	VV	IMP	113	4.8	27.6	27.1
2008/11/14	00:46:23	HH	IMP	23	6.1	24.6	21.5
2008/11/17	00:51:57	VV	IMP	354	6.5	21.4	17.5
2008/12/22	00:51:59	VV	IMP	91	2.0	22.6	17.9

**Appendix 2.** Same as Appendix 1, but for 73 images and *in situ* measurements at Shirahama.

SAR				<i>In situ</i>		Sea Temp. ( °C)	Air Temp. ( °C)
Date (year/month/day)	Time (hr:min:sec)	Polarization	Observation Mode	Wind Direction ( °C)	Wind Speed (m/s) at 10 m		
2003/05/07	01:09:47	VV	IMP	114	3.6	22.1	22.4
2003/07/16	01:09:53	VV	IMP	296	6.6	25.2	23.6
2003/09/24	01:09:56	VV	IMP	59	11.2	26.2	22.4
2003/10/29	01:09:50	VV	IMP	347	10.3	24.2	16.8
2003/12/19	01:07:03	VV	IMP	347	15.1	20.2	8.3
2004/01/23	01:07:00	VV	IMP	322	7.8	16.9	4.0
2004/02/11	01:09:51	VV	IMP	320	3.9	17.0	8.3
2004/02/27	01:07:00	VV	IMP	308	9.9	16.4	6.9
2004/05/07	01:07:00	VV	IMP	268	4.5	20.5	20.7
2004/06/30	01:09:55	VV	IMP	64	5.3	25.6	24.7



## Appendix 2. Cont.

SAR				<i>In situ</i>			
Date	Time	Polarization	Observation	Wind	Wind Speed (m/s)	Sea Temp.	Air Temp.
(year/month/day)	(hr:min:sec)		Mode	Direction ( °C)	at 10 m	( °C)	( °C)
2004/07/31	12:48:26	VV	IMP	115	11.3	27.0	28.6
2004/08/20	01:07:04	VV	IMP	259	5.1	27.7	27.6
2004/09/08	01:09:55	VV	IMP	286	2.5	27.2	26.1
2004/10/13	01:09:56	VV	IMP	20	8.9	24.8	21.8
2004/10/29	01:07:06	VV	IMP	93	5.7	22.8	19.6
2005/01/07	01:06:58	VV	IMP	10	7.3	16.2	10.7
2005/02/11	01:07:01	VV	IMP	331	6.1	15.8	5.8
2005/05/11	01:09:59	VV	IMP	247	3.5	17.6	18.2
2005/05/27	01:07:07	VV	IMP	176	3.8	19.0	23.1
2005/10/14	01:07:05	VV	IMP	96	5.0	25.7	24.7
2005/11/18	01:07:03	VV	IMP	6	6.8	22.0	13.2
2005/12/23	01:06:57	VV	IMP	326	9.6	15.4	6.8
2006/01/11	01:09:42	VV	IMP	358	5.9	18.2	8.5
2006/03/03	01:06:54	VV	IMP	317	9.1	15.1	5.9
2007/08/29	01:09:47	VV	IMP	244	2.4	29.1	28.5
2007/11/07	01:09:43	VV	IMP	16	4.0	23.9	18.0
2007/11/23	01:06:48	VV	IMP	352	9.9	21.8	10.3
2007/12/08	12:48:10	VV	IMP	176	5.7	20.5	11.1
2007/12/09	01:03:59	VV	IMP	340	8.3	20.4	10.4
2008/01/12	12:48:12	VV	IMP	348	8.8	20.3	9.8
2008/01/13	01:04:01	VV	IMP	347	9.2	20.0	6.5
2008/01/16	01:09:43	VV	IMP	349	7.1	18.4	7.9
2008/01/31	12:51:01	VV	IMP	7	6.6	17.6	6.0
2008/02/01	01:06:50	VV	IMP	355	5.5	17.5	5.7
2008/02/16	12:48:09	VV	IMP	323	10.8	15.3	4.7
2008/02/17	01:03:59	VV	IMP	342	10.6	15.1	1.7
2008/02/20	01:09:42	VV	IMP	345	6.6	16.8	6.9
2008/03/06	12:51:02	VV	IMP	72	4.9	16.7	7.0
2008/03/07	01:06:51	VV	IMP	337	6.7	16.6	9.0
2008/03/22	12:48:13	VV	IMP	244	4.2	17.5	12.3
2008/03/23	01:04:02	VV	IMP	159	7.3	18.3	13.5
2008/03/26	01:09:43	VV	IMP	354	5.1	18.2	12.9
2010/06/25	01:01:06	VV	WSM	84	3.1	24.1	23.4
2010/07/11	01:01:03	VV	WSM	168	7.4	24.0	26.1
2010/07/27	01:01:00	VV	WSM	184	2	27.0	30.7
2010/07/30	01:01:06	VV	WSM	205	2.9	26.4	27.8
2010/08/12	00:00:57	VV	WSM	218	10.6	27.8	28.2
2010/08/15	01:01:03	VV	WSM	211	5.2	28.4	28.8
2010/08/18	01:01:09	VV	WSM	286	4.2	28.2	28.7
2010/08/31	01:01:00	VV	WSM	223	2.4	29.0	29.3
2010/09/03	01:01:06	VV	WSM	267	3	29.6	28.3

## Appendix 2. Cont.

SAR				In situ			
Date	Time	Polarization	Observation	Wind	Wind Speed (m/s)	Sea Temp.	Air Temp.
(year/month/day)	(hr:min:sec)		Mode	Direction ( °C)	at 10 m	( °C)	( °C)
2011/10/30	01:01:07	VV	WSM	30	2.8	24.2	19.8
2011/12/07	01:01:14	VV	WSM	359	5.7	21.7	14.5
2011/12/10	01:01:04	VV	WSM	33	4.3	21.5	9.5
2011/12/18	01:01:11	VV	WSM	3	9.1	21.2	8.4
2010/06/27	12:12:56	VV	WSM	187	5	24.0	25.4
2010/07/13	12:12:53	VV	WSM	164	7.6	23.9	26.1
2010/08/14	12:12:47	VV	WSM	176	5.6	28.2	28.0
2010/09/05	12:12:56	VV	WSM	104	2.7	29.9	27.8
2010/09/18	12:12:47	VV	WSM	95	3.2	27.7	24.0
2010/09/21	12:12:53	VV	WSM	103	2.9	27.7	26.4
2011/10/18	12:12:58	VV	WSM	108	3.9	25.4	18.5
2011/10/26	13:13:04	VV	WSM	29	5.2	24.5	14.6
2011/10/29	12:12:54	VV	WSM	119	5.1	24.2	19.1
2011/11/06	13:13:01	VV	WSM	355	5.8	23.8	21.5
2011/11/25	13:13:04	VV	WSM	18	5.9	21.8	11.0
2011/12/06	13:13:01	VV	WSM	23	4.7	21.8	13.6
2011/12/17	12:12:58	VV	WSM	26	5.7	20.3	7.9
2011/12/28	12:12:55	VV	WSM	71	3.2	18.9	6.5
2012/01/06	01:01:15	VV	WSM	14	7.3	18.0	6.6
2012/01/09	01:01:05	VV	WSM	24	7.1	17.9	9.5
2012/01/05	13:13:01	VV	WSM	27	7.2	18.3	6.3
2012/01/16	12:12:58	VV	WSM	46	7	18.3	6.4

## Appendix 3. The model configuration used in the WRF simulation is shown.

Initial data	MANAL NGSST (0.05 ° × 0.05 °, daily)		
Vertical resolution	28 levels (surface to 10 hPa)		
Nesting option	two-way nesting		
Domain	Domain 1	Domain 2	Domain 3
Horizontal	4.5 km	1.5 km	0.5 km
Grid points	100 × 100	100 × 100	100 × 100
Time step	27 sec	9 sec	3 sec
Physics option	Surface layer	Monin-Obukhov (Janjic Eta)	
	Planetary Boundary Layer	Mellor-Yamada-Janjic (Eta) TKE	
	Short wave radiation	Dudhia	
	Long wave radiation	RRTM	
	Cloud microphysics	WSM3	
	Cumulus parameterization	Kain-Fritsch (new Eta)	none
	Land surface	Five-layer soil	

**Appendix 3. Cont.**

4dda option	Enable including PBL	Enable excluding PBL	Enable excluding BL
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MANAL: Japan Meteorological Agency Meso-Analysis; NGSST: New Generation Sea Surface Temperature;  
PBL: Planetary Boundary Layer.

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